

Ocean-atmosphere coupling with OASIS

Large scale meteorology and climate research group

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CNRM/GAME

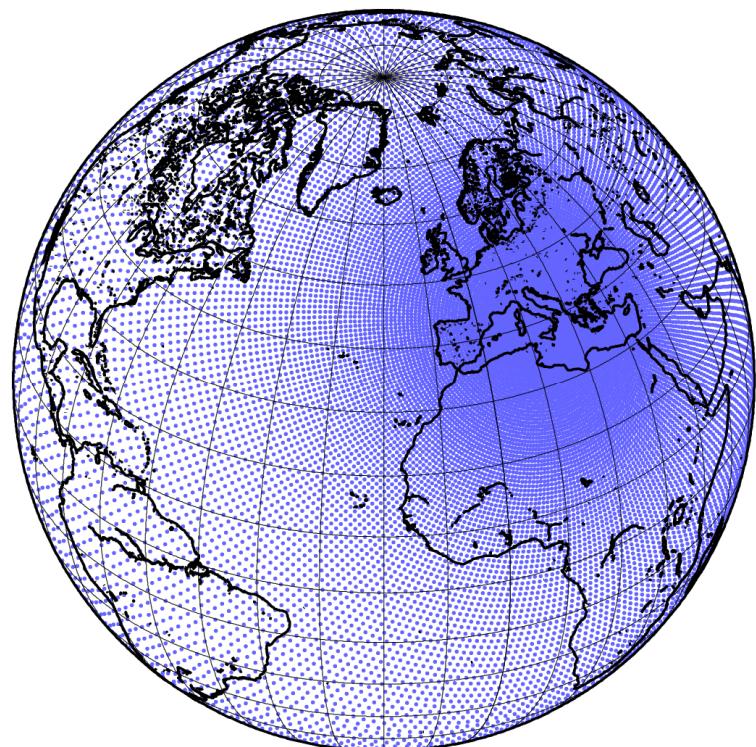
Ocean-atmosphere coupling with OASIS

- Regional coupling on the Mediterranean region
- Conservation issue of the non-solar flux from atmosphere to ocean

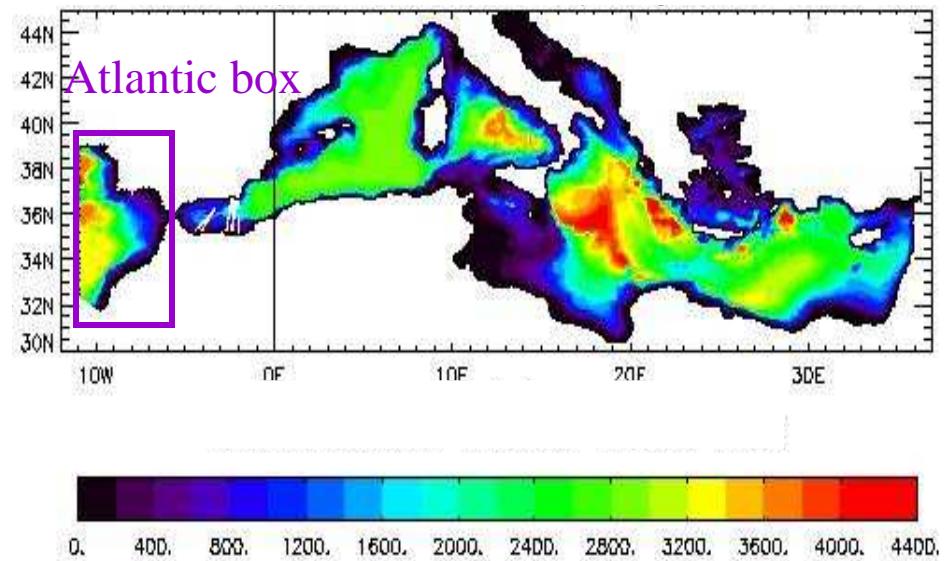
Regional coupling on the Mediterranean region (I)

ARPEGE-Climate v3 :

50 km resolution on the Mediterranean



NEMOMED8: a NEMO-V2 regional version
(LOCEAN); about 10 km resolution

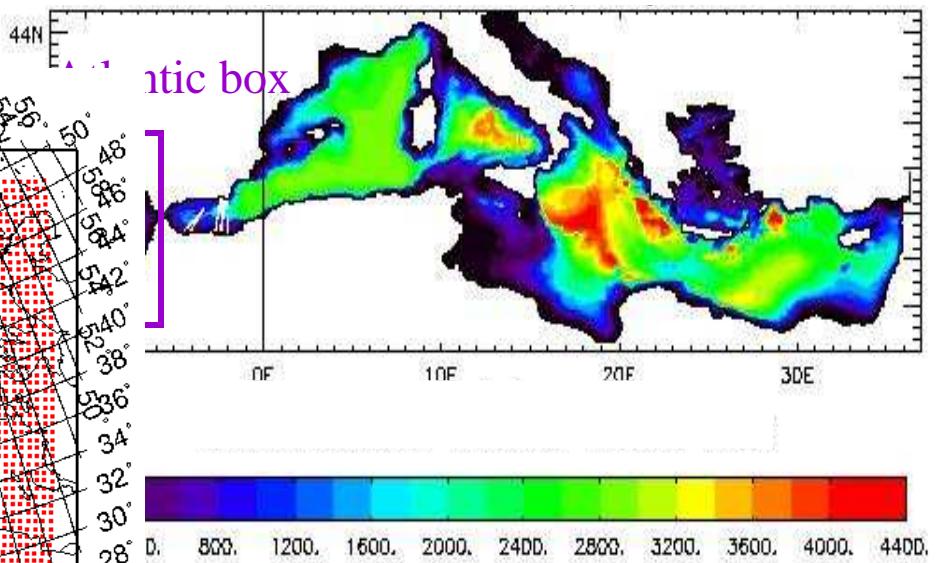
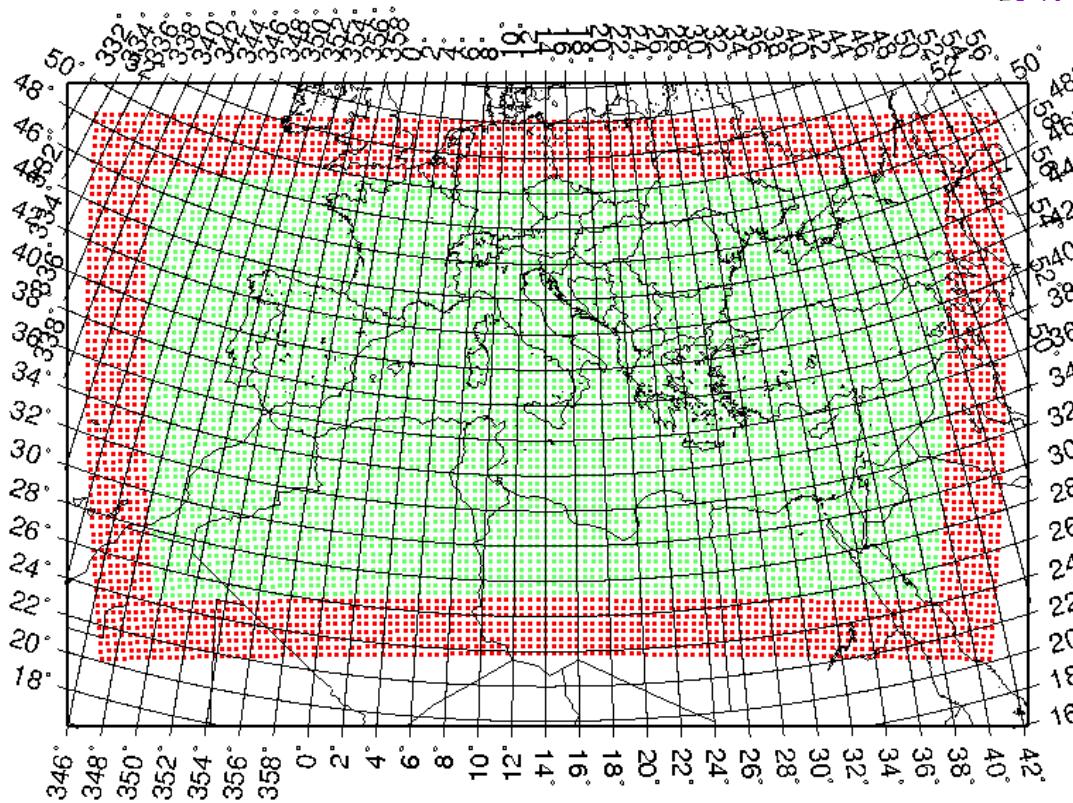


Regional coupling on the Mediterranean region (I)

ALADIN-Climate:

50 km resolution on the Mediterranean

NEMOMED8: a NEMO-V2 regional version
(LOCEAN); about 10 km resolution



Regional coupling on the Mediterranean region (I)

Scenarios of climate change: comparison regional-global atm-ocean coupling

Already done with ARPEGE-Climate – OASIS2.4 – OPAMED8: 21st century climate change scenario for the Mediterranean using a coupled atmosphere-ocean regional climate model, Somot et al., Global and Planetary Change, 2008.

Characteristics of the coupling (ARPEGE-Climate):

- OASIS3
- ARPEGE: global 1D ($n,1$) grid stretched, center in the Tyrrhenian Sea
- Oceanic grid: tilted at the Gibraltar strait (not regular)
- MASK-EXTRAP-MOZAIC or INTERP from atm to ocean
- MOZAIC-FILLING from ocean to atm: the ocean fields on the Mediterranean are filled with global data (observations or large scale model outputs previously interpolated on the atm grid)

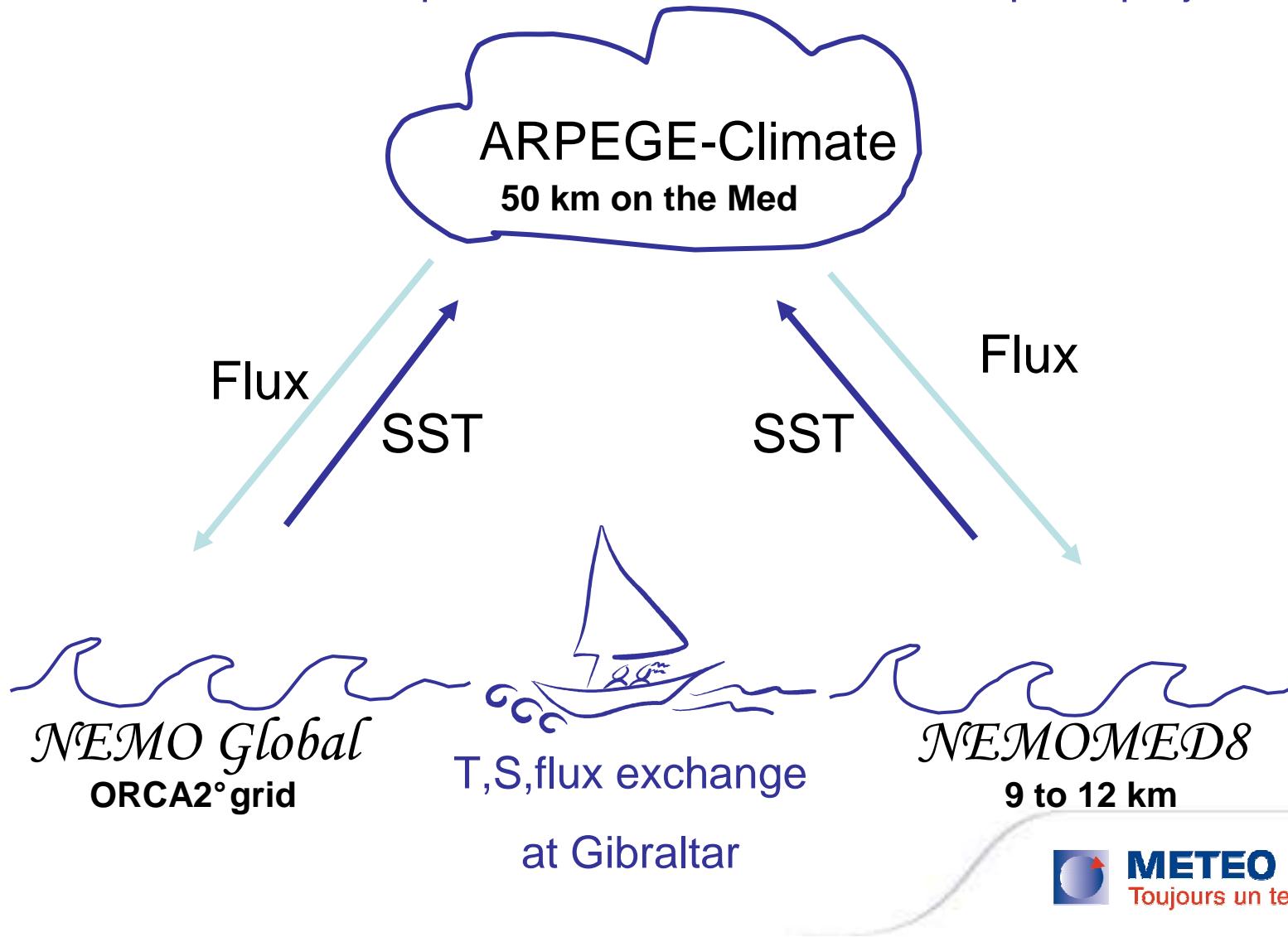
Characteristics of the coupling (ALADIN-Climate):

- OASIS3
- ALADIN 2D grid
- MASK-EXTRAP-SCRIPR for atm to ocean
- SCRIPR-FILLING from ocean to atm

For both systems, no problem of conservation of the fluxes during the interpolation.

Regional coupling on the Mediterranean region (II)

The tri-coupled model of the CIRCE european project



Regional coupling on the Mediterranean region (II)

Characteristics of the coupling:

- OASIS3
- ARPEGE sends its fluxes twice, for the two oceanic models
- The SST of the two oceanic models is combined in OASIS:

Regional coupling on the Mediterranean region (II)

```
#####
#  
# Field 2 : sea surface temperature global  
#  
SOSSTSST SISUTES1 1 06400 6 ocglo.nc AUXILIARY  
182 149 35718 1 tnem mdh2 LAG=+4800  
P 2 R 0  
#  
CHECKIN MASK EXTRAP SCRIPR MASKP CHECKOUT  
#  
INT=1  
99999999.  
NINENN 2 0 1  
GAUSWGT LR SCALAR LATON 10 4 2.  
0  
INT=1  
#  
#####
...  
#####
#  
# Field 9 : sea surface temperature tom8  
#  
SOSSTMED SISUTESU 1 06400 4 ocmed.nc EXPORTED  
394 160 35718 1 tom8 medh LAG=+1200  
P 0 R 0  
#  
CHECKIN MOZAIC BLASNEW CHECKOUT  
#  
INT=1  
runoffh2 93 1 38  
1. 1  
SISUTES1 1  
INT=1  
#
```

tnem: ORCA grid

mdh2: ARPEGE grid, mask on NEMOMED8 domain

MASKP: value=0 on the land and on the NEMOMED8 domain

tom8: NEMOMED8 grid

medh: ARPEGE grid

SISUTESU=SISUTESU+SISUTES1

Regional coupling on the Mediterranean region (II)

The communication between the two oceanic model is made by files.

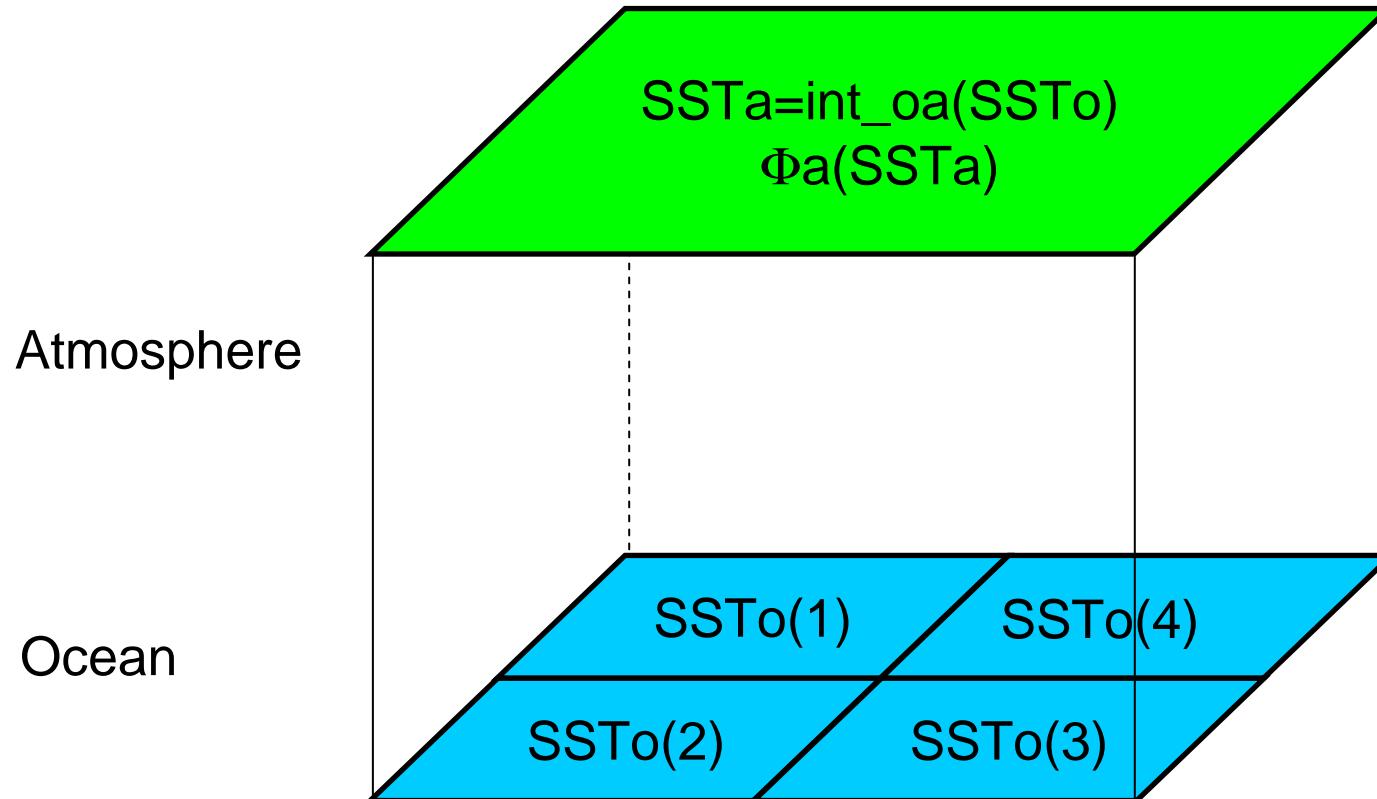
At the Gibraltar Strait, daily:

- Incoming flux Atl-ORCA2 => NEMOMED8: the T and S profiles of NEMO-ORCA2 on the Atlantic box of NEMOMED8 are used for the 3D damping fields
- Outcoming flux NEMOMED8 => Atl-ORCA2 : used in the Cross Land Advection subroutine of NEMO

Each model reads the file written by the other one, no matter if there is a one day shift.

For CIRCE a 100 year A1B scenario is now running.

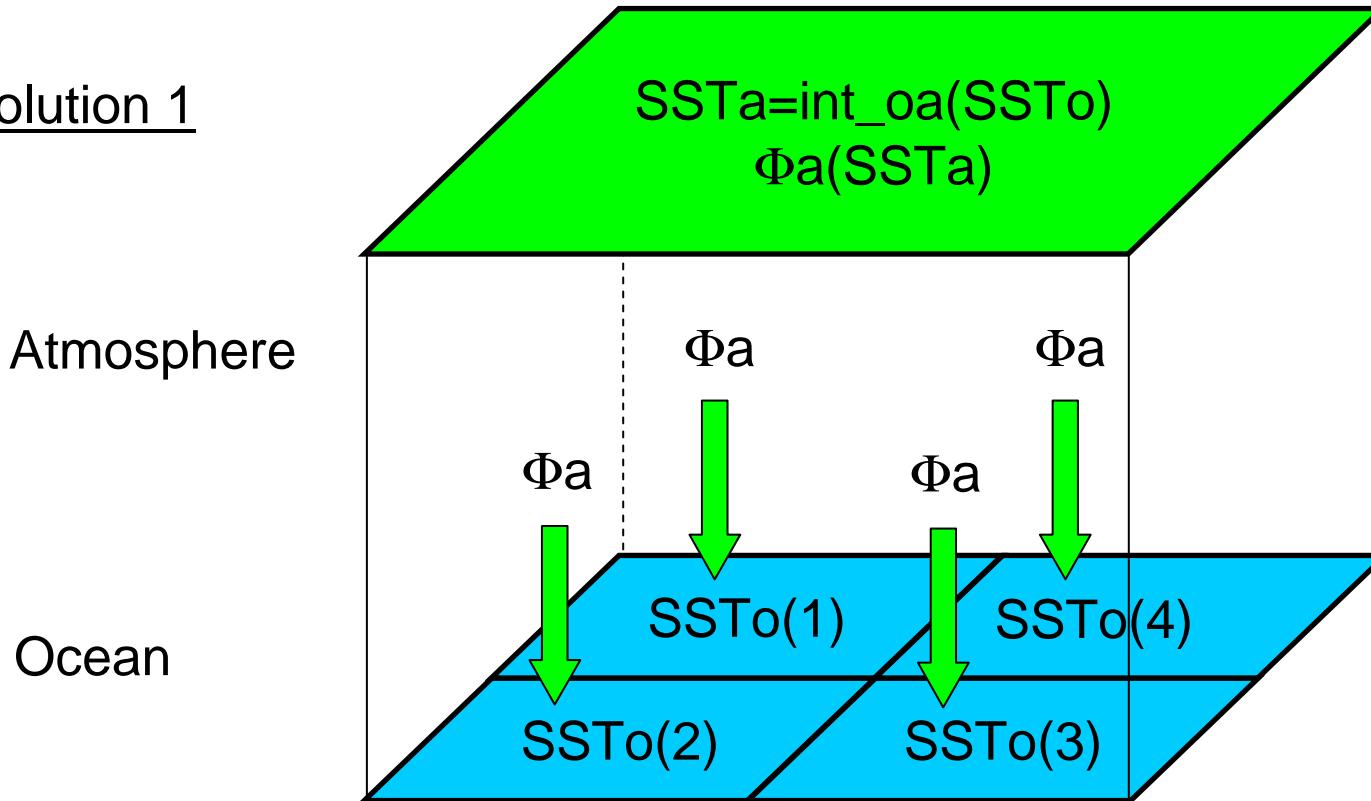
Ocean-atmosphere coupling with OASIS Conservation at the atm-ocean interface



$$\sum_i \alpha(i) \cdot SST_o(i) = SST_a$$

Ocean-atmosphere coupling with OASIS Conservation at the atm-ocean interface

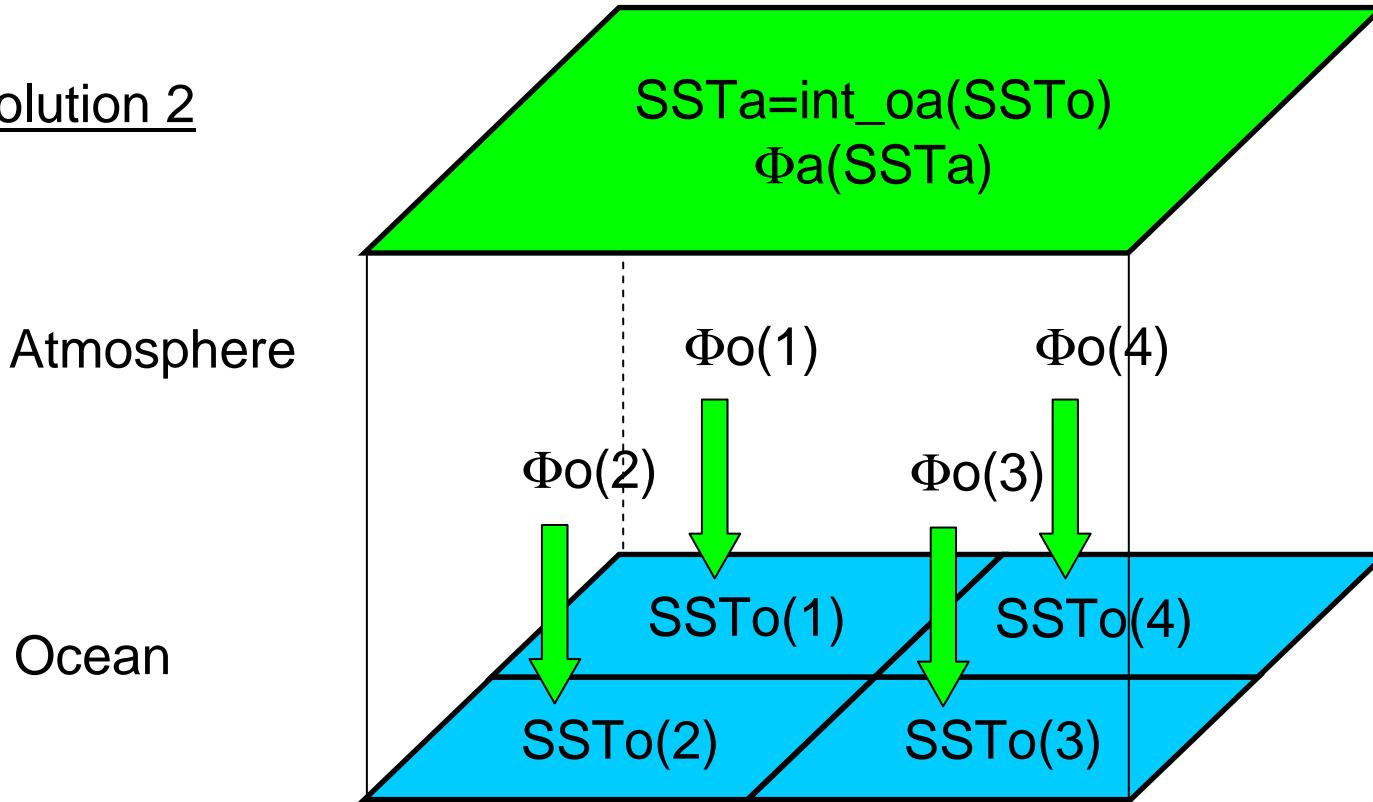
Solution 1



$$\sum_i \alpha(i) \cdot SST_o(i) = SST_a \quad \text{avec} \quad \sum_i \alpha_i = 1 \quad \Phi_o(i) = \Phi_a$$

Ocean-atmosphere coupling with OASIS Conservation at the atm-ocean interface

Solution 2



$$\sum_i \alpha(i) \cdot SST_o(i) = SST_a \quad \Phi_o(i) = \Phi_a + \frac{\partial \Phi_a}{\partial SST_a} (SST_o(i) - SST_a)$$

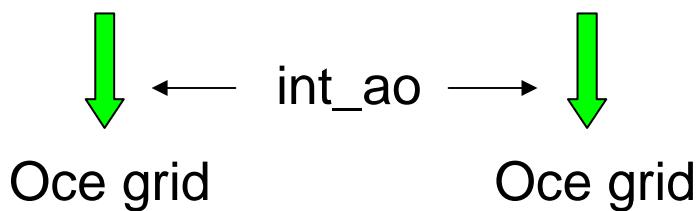
SUBGRID option of OASIS for the non-solar flux

Ocean-atmosphere coupling with OASIS Conservation at the atm-ocean interface

$$\Phi_o(i) = \Phi_a + \frac{\partial \Phi_a}{\partial SST_a} (SST_o(i) - SST_a) \xrightarrow{\times \alpha_i} \sum_i \alpha(i) \cdot \Phi_o(i) = \Phi_a$$

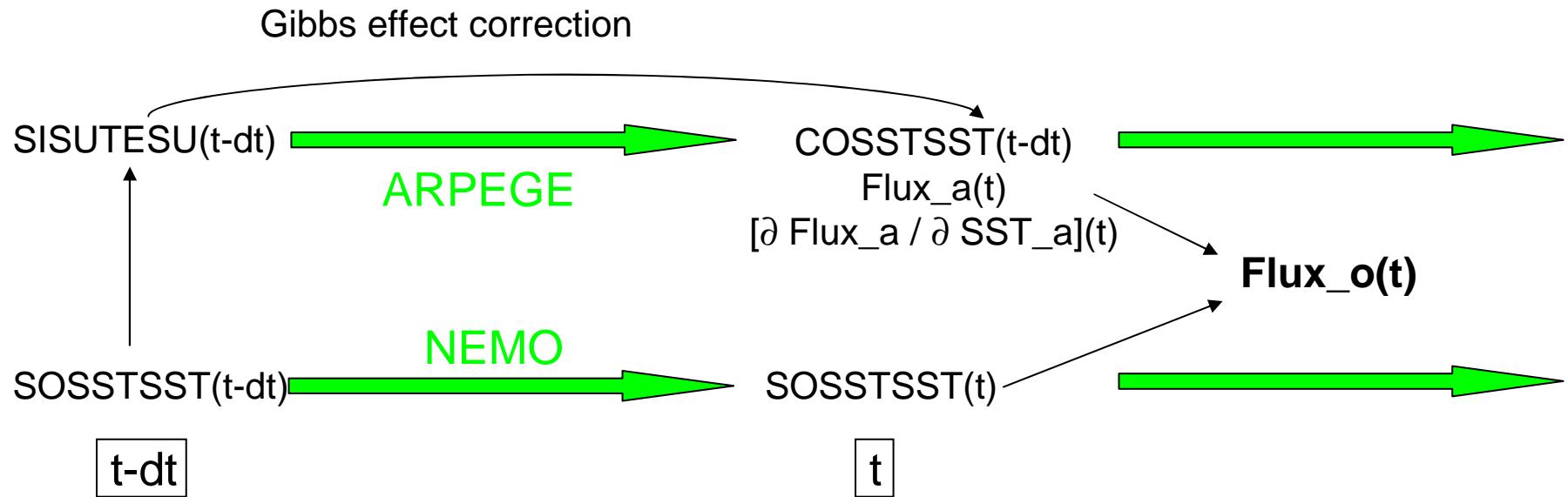
(conservation of the energy)

Oce grid Atm grid Oce grid Atm grid



$$\Phi_o^t(i) = \text{int_ao} \left[\Phi_a^t + \frac{\partial \Phi_a^t}{\partial SST_a} \right] \times [SST_o^t(i) - \text{int_ao}(SST_a^t)]$$

Ocean-atmosphere coupling with OASIS Conservation at the atm-ocean interface



$$\Phi_o^t(i) = \text{int_ao} \left[\Phi_a^t + \frac{\partial \Phi_a}{\partial SST_a} \right] \times [SST_o^t(i) - \text{int_ao}(\text{int_oa}(SST_o^{t-dt}(i)))]$$

« SOSSTSST » « COSSTSST »

Ocean-atmosphere coupling with OASIS Conservation at the atm-ocean interface

In summary, the use of SUBGRID gives:

$$\Phi_o^t(i) = \text{int_ao} \left[\Phi_a^t + \frac{\partial \Phi_a}{\partial SST_a}^t \right] \times [SST_o^t(i) - \text{int_ao}(\text{int_oa}(SST_o^{t-dt}(i)))]$$

- the fields are not simultaneous
- the correction of the Gibbs effect leads to other errors

→ $\sum_i \alpha(i) \cdot \Phi_o(i) = \Phi_a$ The non-solar flux is not conserved (loss ~1%)

→ The use of CONSERV (OASIS) is a bad solution: in CNRM-CM3 loss~2,5%

Ocean-atmosphere coupling with OASIS Conservation at the atm-ocean interface

One solution is to modify « manually » the weights used in OASIS for the interpolation, to combine int_ao and int_oa in one.

$$\Phi_o^t(i) = \text{int_ao} \left[\Phi_a^t + \frac{\partial \Phi_a}{\partial SST_a}^t \right] \times [SST_o^t(i) - \text{int_ao}(\text{int_oa}(SST_o^{t-dt}(i)))]$$


$$[SST_o^t(i) - \text{int_ao}(\text{int_oa}(SST_o^t(i)))]$$

→ No more time shift

Loss ~ 0,01W.m⁻²



Thank you.